THERMAL ANALYSIS OF PALYGORSKITE

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ABSTRACT

Slickensides and fissures in a fresh outcrop of silicified limestones at the Bürgenstock (Switzerland) contained palygorskite. According to TG, DTA and heating X-ray experiments, dehydration of the palygorskite took place in four steps. Each step clearly showed an alteration in the unit cell dimensions. Exothermic reactions which occur at higher temperatures resulted in the formation of a Mg-Al-silicate high phase together with some clinoenstatite and the formation of β -cristobalite with some cordierite, respectively.

INTRODUCTION

Müller-Vonmoos and Schindler¹ described palygorskite on slickensides and in fissures of silicified limestones at the Bürgenstock (Switzerland). The palygorskite formed dirty white to yellowish coatings (Fig. 1). In thin sections the mineral could

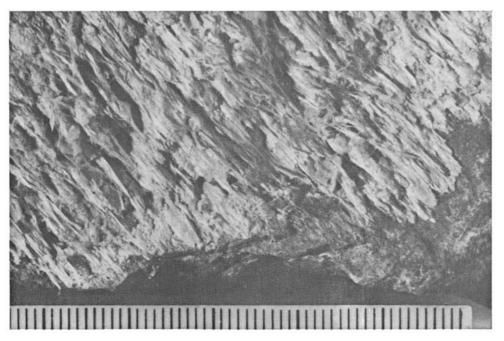


Fig. 1. Palygorskite forming coatings on silicified limestone (scale divisions 1 mm).



Fig. 2. Palygorskite in a fissure of silicified limestone (thin section, 80×).

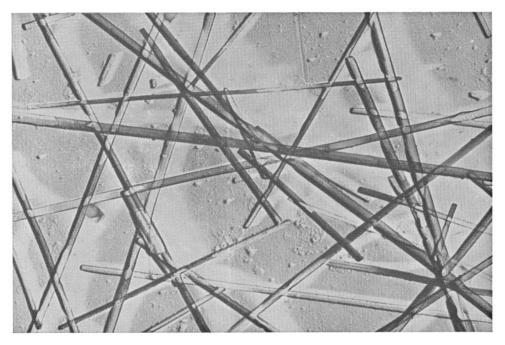


Fig. 3. Electron micrograph of palygorskite $(22500 \times)$.

be observed in fissures, mostly containing calcite (Fig. 2). Electron micrographs revealed the typical lath-shaped units (Fig. 3). The rock belongs to Lower Cretaceous (Hauterivian) of a non-metamorphic helvetic nap. Probably the palygorskite grew during the last orogenic phase (Pliocene) in low temperature solutions. According to chemical analysis and thermogravimetric investigations, the following composition can be given:

 $(Si_{7.88}Al_{0.12})_{8.00}(Al_{1.55}Mg_{1.92}Fe_{0.46})_{3.93}(OH)_{1.95}(OH_2)_{4.13} \cdot 3.90H_2O, Ca_{0.06}K_{0.03}Na_{0.10}$

Thermal analysis (TG-DTA) of this palygorskite indicated four endothermic reactions and corresponding weight loss steps and three exothermic reactions. The interpretation of these reactions given by Müller-Vonmoos and Schindler¹ could be proved by X-ray investigations.

SAMPLE PREPARATION AND EXPERIMENTAL TECHNIQUE

For TG and X-ray investigations the palygorskite was purified from calcite and quartz by hand picking. For the simultaneous TG-DTA the calcite was removed by boiling in an acetate-buffer (15 min, pH 5) followed by changing the palygorskite into the Ca-form.

TG and TG-DTA were carried out with a Mettler-thermobalance² which permitted simultaneous recording of weight changes, DTA-curves and the actual temperature during the test. The sample weight was 50 mg. For TG and TG-DTA we

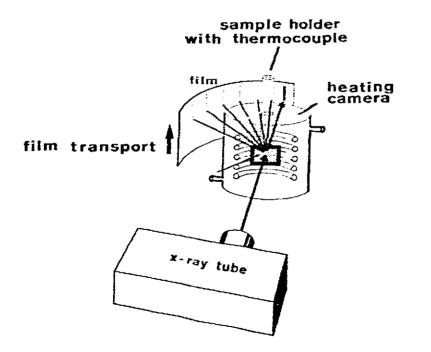


Fig. 4. Schematic sketch of X-ray high temperature camera and X-ray tube.

heated in flowing air at heating rates of 0.5° C min⁻¹ and 6° C min⁻¹, respectively. The reference material in DTA-runs was dead burnt palygorskite.

An Enraf Nonius high temperature X-ray camera (system Lenne-de Wolff) was used to follow the decomposition of palygorskite. This method has the advantage that the sample can be mounted in the sample holder of the X-ray camera and be heated up continuously during exposure of the film without changing its position (Fig. 4). The X-ray camera is equipped with a quartz monochromator. A special film holder allows a transport of the film at various rates. Time marks are printed automatically on the film for correlating the X-ray patterns to the specific times in the simultaneously recorded temperature curve³. The palygorskite was heated in flowing air at a heating rate of 0.5° C min⁻¹.

RESULTS AND DISCUSSION

Heating the palygorskite at 0.5° C min⁻¹, the dehydration took place in three steps (Fig. 6). If it is assumed that the three steps correspond to the loss of zeolithic water, bonded water and the dehydroxylation, respectively, the value obtained for the bonded water is too low, whereas the OH-content turns out to be too high (Table 1). However, when heating at a rate of 6°C min⁻¹, the third step of weight loss clearly separated into two steps and the simultaneous DTA indicated two endothermic reactions, at 525 and 598°C. Assuming the dehydration of the bonded water in two steps (TG 0.5° C min⁻¹, 85–195°C and 195–~480°C), the loss of bonded water and the dehydroxylation amounts to 8.8 and 2.1%, respectively.

TABLE I

DEHYDRATION OF PALYGORSKITE HEATED AT 0.5° C min⁻¹ (TG) AND 6^oC min⁻¹ (TG-DTA), RESPECTIVELY, AND THE VALUES FOR THE IDEAL UNIT CELL Si₈Mg₅O₂₀(OH)₂(OH)₂(OH)₂·4H₂O AFTER BRADLEY⁸

Water content of the ideal unit cell		ΤG		TG-DTA		
Zeolithic water	8.6%	up to 85°C	7.9%	up to 195°C	8.5%	
Bonded water	8.6%	85-195°C	4.3%	195–322 °C 322–530 °C	4.7% 4.3%	
Hydroxyl water	2.15%	195–640 °C	6.6%	530-820°C	2.1%	
Total	19.35%		18.8%		19.6%	

When heating at a rate of $0.5 \,^{\circ}$ C min⁻¹, the loss of zeolithic water took place at temperatures up to $85 \,^{\circ}$ C (Fig. 6). By X-ray between 35 and 55 $\,^{\circ}$ C, a distinct change of the lattice constants, especially a contraction in the direction of the a-axis, could be observed (Fig. 5, Table 2). According to this unit cell deformation, the zeolithic water should be released completely at 55 $\,^{\circ}$ C. On the other hand, the loss of the zeolithic water observed by TG was completed only at $85 \,^{\circ}$ C. An explanation for this delay may be a readsorption of the zeolithic water on the external surface. A loss of zeo-

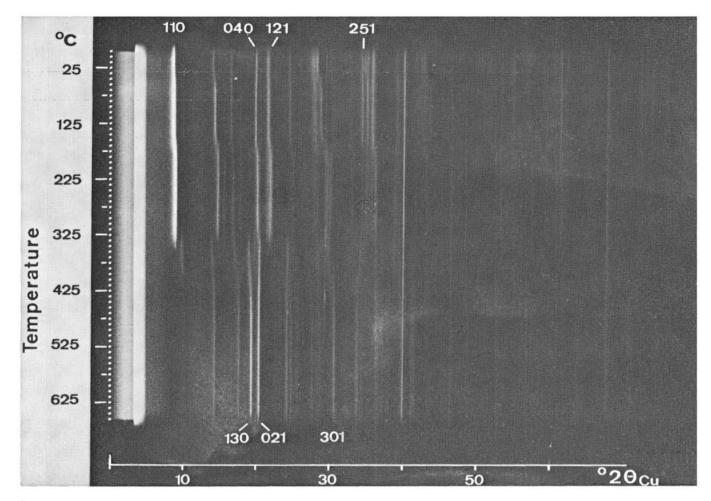


Fig. 5. X-ray film showing the thermal dehydration of palygorskite. $Cu-K_3$ radiation; film speed, 5 mm h⁻¹; heating rate, 0.5 °C min⁻¹.

TABLE 2

d-VALUES, ESTIMATED INTENSITIES AND INDICES OF REFLECTIONS OF PALYGORSKITE AT 25, 55, 195 AND 475°C Copper- K_x -radiation, heating rate 0.5°C min⁻¹. Impurities: C = calcite. M = muscovite.

25°C		55°C		195°C			475°C				
d	I	hkl	d	1	hkl	đ	I	hki	đ	1	hkl
10.254	100	110	10.178	100	110	9.889	100	110	8.640	10	110
6.302	40	200	6.204	50	200	6.001	40	200	6.186	60	120
						5.679	20	210	5.023	50	011
5.371	30	130	5.368	30	130	5.268	30	130	4.583	100	130
			4.986	<10	M	4.976	20	001/M	4.354	100	021

TABLE 2 (continued)

25°C			55 °C			195° C			475°C		
đ	Ι	hkl	đ	I	hki		I	hki		I	hki
4.451	60	040	4. 46 2	60	040	4.395	60	040	3.674	40	031
4.353	10	121	4.34I	10	121				3.590	20	140
						4.340	20	021			
									3.466	10	131
						4.127	30	140	2 2 2 2		•
4.126	60	121	4.116	60	121	4.071	50	121	3.223	10	?
4.120	60	121	4.110	00	121	4.071	20	121	3.135	10	041
						4.012	20	300	5.155	10	V+1
									3.04	10	С
3.637	30	240	3.627	30	240	3.555	20	240			
									2.916	40	301
3.45	<10	150									_
									2.651	30	051
3.221	50	231	3.198	40	231				2.443	10	122
						3.166	10	141	£44J	10	122
						5.100	10	141	2.358	<10	032
						3.029	20	250/C			
									2.266	10	222
3.158	60	400	3.102	40	400	3.000	20	400			
									2.181	20	042
3-028	10	С	3.025	10	С						
						2 067	20	410/740	1.723	10	264/540
2_604	40	251	2.599	40	251	2.957 2.598	30 20	410/340 251			
2004		201	2	40	<u>.</u>	2.558	20	411			
2.576	40	440/061				2.480	40	440			
2.551	40	161	2.555	40	161						
2.502	60	161	2.505	50	161						
			2.295	10	212						
			2.280	10	232						
			2.242	10	222						
2.122		271?	3 107	10	200						
2.104	20	600	2.107	10 10	280 600						
2.104	20	000	2.067	10	000	1.946	10	620			

Dimensions of the elementary cells

ΜοποςΙίπις	Monoclinic	Orthorhombic	Orthorhombic
a = 12.65 Å b = 17.80 Å c = 5.22 Å β = 95°30'	a = 12.43 Å b = 17.86 Å c = 5.21 Å β = 95°33'	a = 12.00 Å b = 17.59 Å c = 4.98 Å	a = 10.45 Å b = 15.28 Å c = 5.32 Å

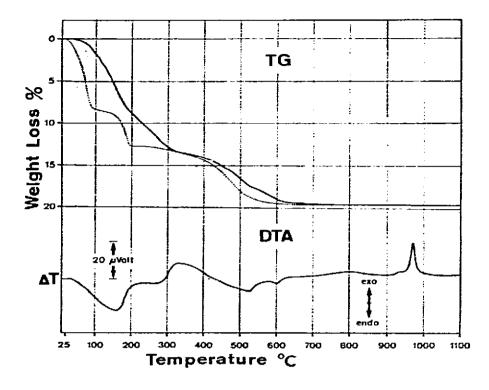


Fig. 6. TG and DTA curves of palygorskite., TG-curve, heating rate 0.5°C min⁻¹, air flow 3 l h⁻¹., TG-DTA curves recorded simultaneously, heating rate 6°C min⁻¹, air flow 6 l h⁻¹.

lithic water in two steps was also observed by Veniale⁴ investigating palygorskite from Attapulgus.

Based on TG (heating rate 0.5° C min⁻¹), the first loss of bonded water took place between 150 and 195°C. In the same temperature range a change could be observed in the X-ray pattern. According to TG, the second loss of bonded water started gradually above 200°C. Between 335 and 450°C, the lattice changed due to the formation of palygorskite-anhydrite. The two steps of dehydration might be attributed to the different bond strength of the water to Mg and Al, respectively. Between 645 and 675°C the anhydrite decomposed into an amorphous phase.

Between 740 and 885°C X-ray photographs indicated the crystallisation of a Mg-Al-silicate phase (Mg₂Al₄Si₅O₁₈) with high quartz structure⁵ besides some clinoenstatite. In the same temperature range DTA revealed a weak exothermic reaction. Additional exothermic peaks at 930 and 975°C corresponded to the formation of β -cristobalite and some cordierite.

According to our investigations, each of the four dehydration steps of the palygorskite were clearly correlated to structural changes. This is not in agreement with the investigations of Hayashi et al.⁶ and Preisinger⁷, where no changes were observed in the structure of palygorskite during dehydration up to 210° and 350°C, respectively.

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